

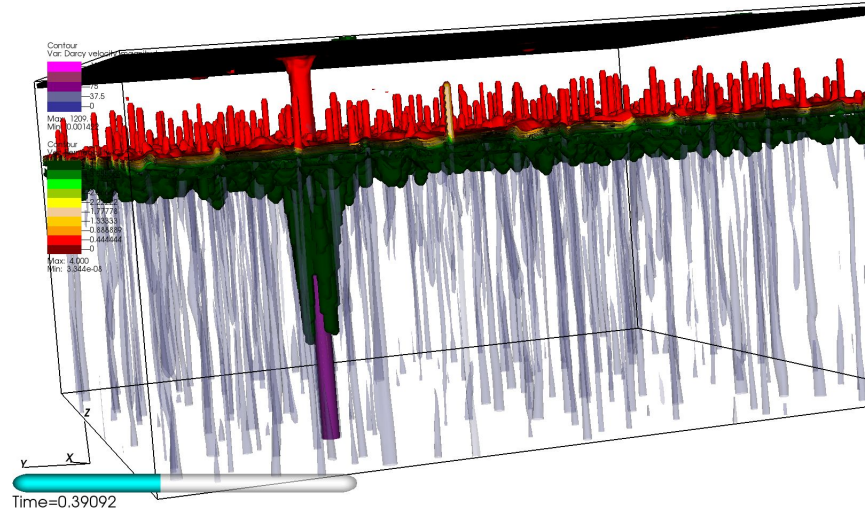
3D convection, phase change, and solute transport in mushy sea ice

Dan Martin, James Parkinson, Andrew Wells, Richard Katz

Lawrence Berkeley National Laboratory (USA), Oxford University (UK).

Summary:

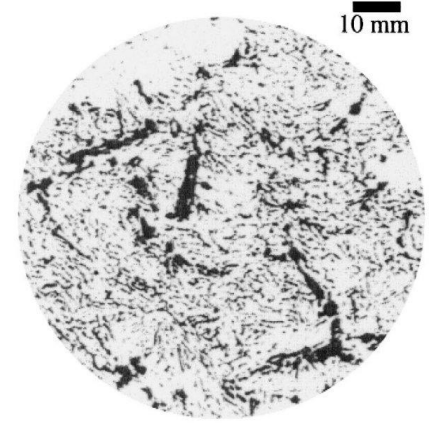
- Simulated brine drainage via 3-D convection in porous mushy sea ice
- Shallow region near ice-ocean interface is desalinated by many small brine channels
- Full-depth “mega-channels” allow drainage from saline layer near top of ice.



What is a mushy layer?

Dense brine drains convectively from porous mushy sea ice into the ocean.

- What is spatial structure of this flow in 3 dimensions?



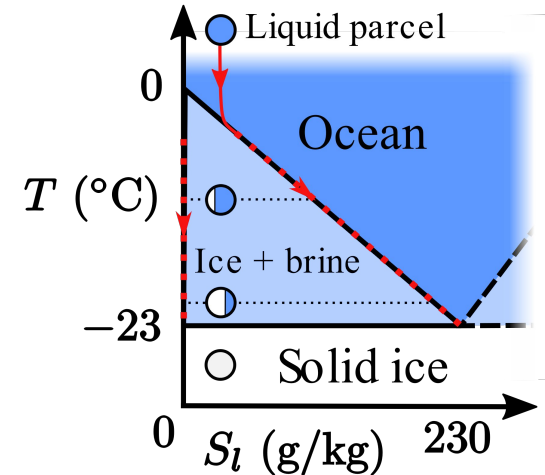
Upper fig.: Sea ice is a porous mixture of solid ice crystals (white) and liquid brine (dark).

H. Eicken et al. Cold Regions Science and Technology 31.3 (2000), pp. 207–225

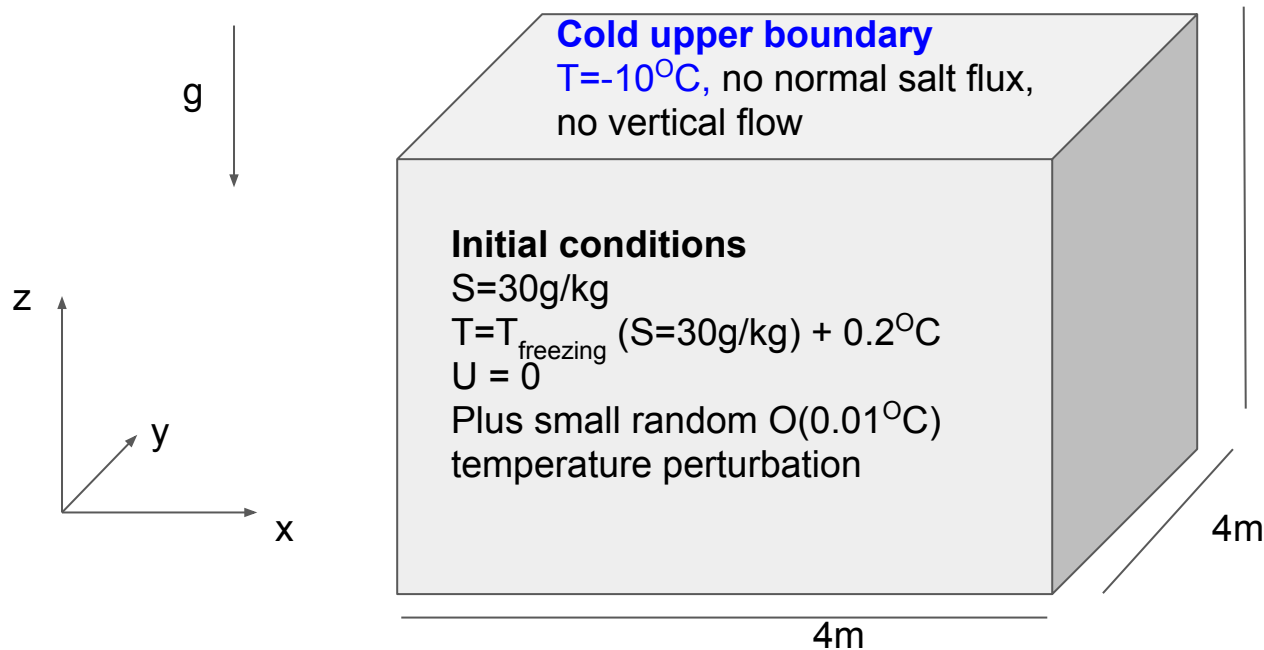
Lower fig.: Trajectory (→) of a solidifying salt water parcel through the phase diagram. As the temperature T decreases, the ice fraction increases and the residual brine salinity S_l increases making the fluid denser, which can drive convection.

Using a linear approximation for the liquidus curve, the freezing point is

$$T_f(S_l) = -0.1S_l.$$



Problem setup



Numerically solve
mushy-layer equations
for porous ice-water
matrix
(see appendix).

Horizontally periodic

Open bottom boundary

Inflow/outflow, with constant pressure

Inflow: $S = 30\text{g/kg}$, $T = T_{\text{freezing}}(S = 30\text{g/kg}) + 0.2^{\circ}\text{C}$

Movie

Contours of:

Ice permeability

-function of ice
porosity;

(red lower,

green higher

~ice-ocean interface)

Velocity

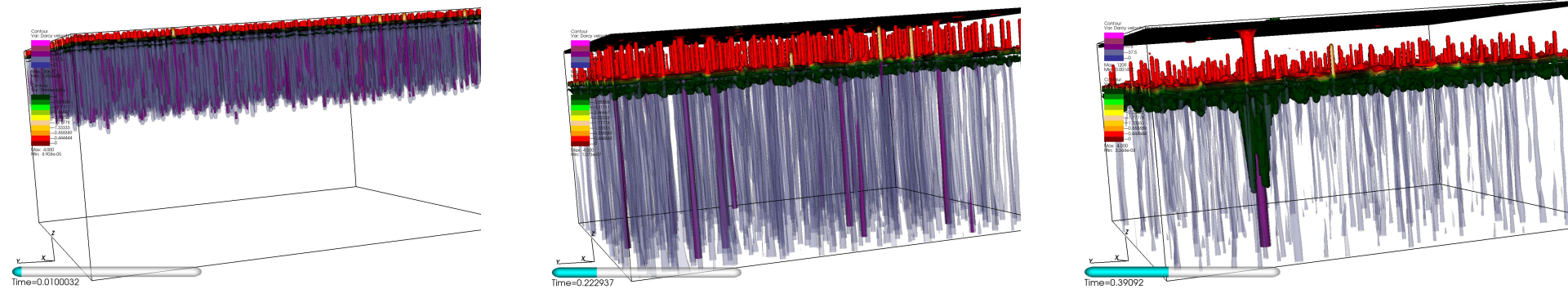
(blue lower,

purple higher).

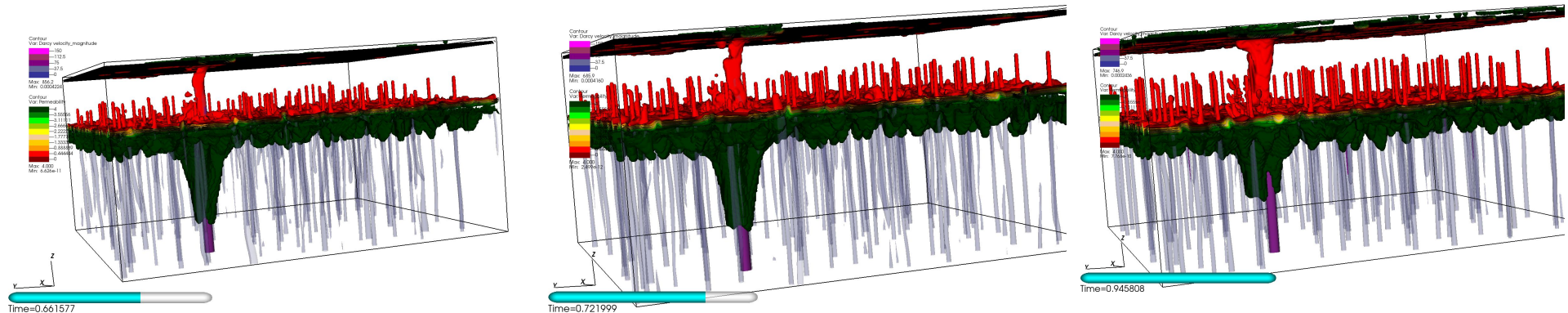
https://drive.google.com/file/d/1JBItmurLZ1zHKXT-Qt-8pmVEHuJIYdKI/view?usp=sharing_eil&ts=5eaef8d6



Results -- Permeability and Velocity



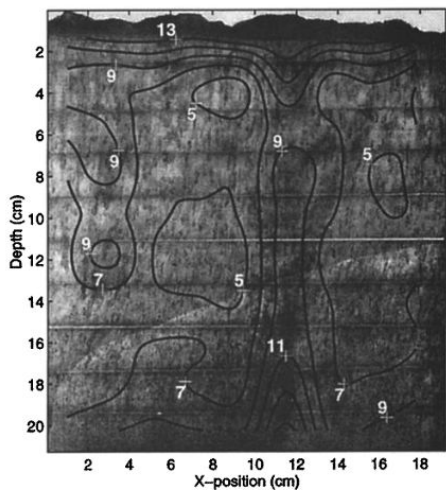
Time (each row) →



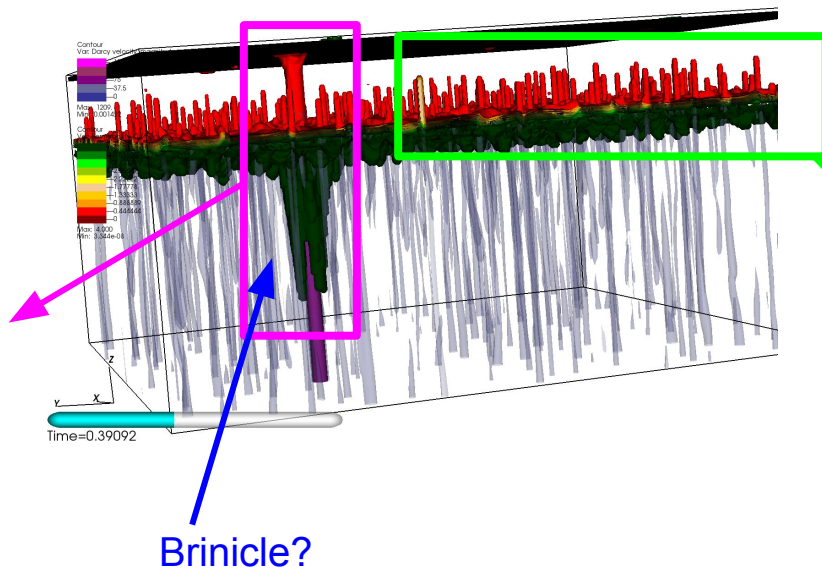
Fine channels coarsen, and mega-channel forms as time progresses.

Qualitative similarities with experiments

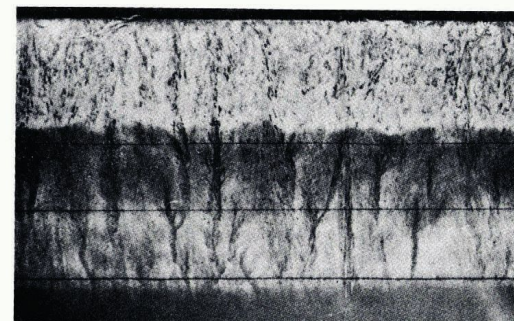
Large “mega-channel”



Contours of bulk salinity (psu)
Fig 6d from Cottier & Wadhams (1999)

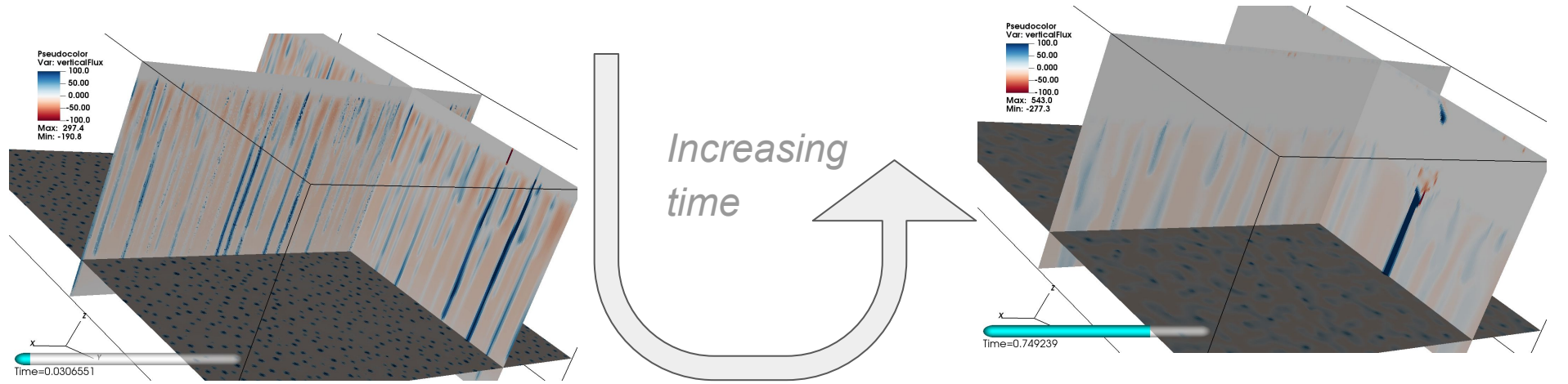


Array of smaller brine channels



Photograph of dye entrainment in sea ice
Fig 3c from Eide & Martin (1975)

Results -- Vertical Salinity Flux

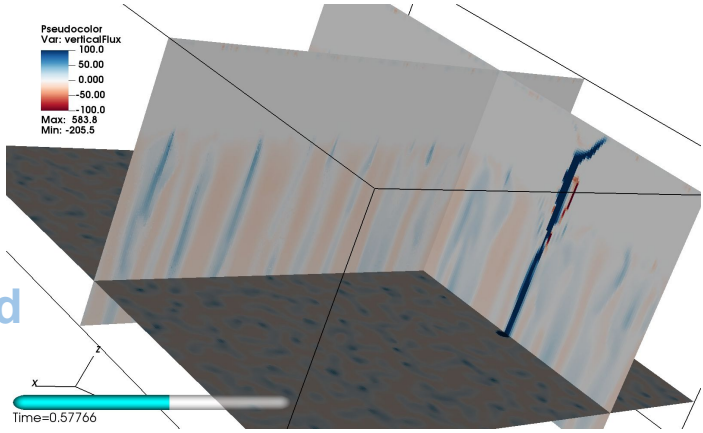


Vertical salt flux:

Dark Blue - Strong Downward

Light Blue - Weak downward

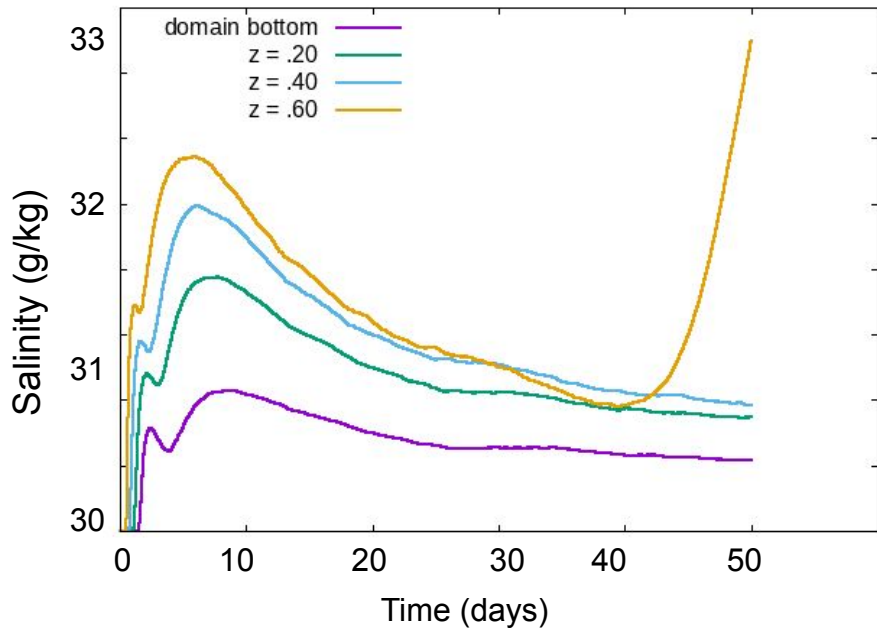
Pink - weak upward.



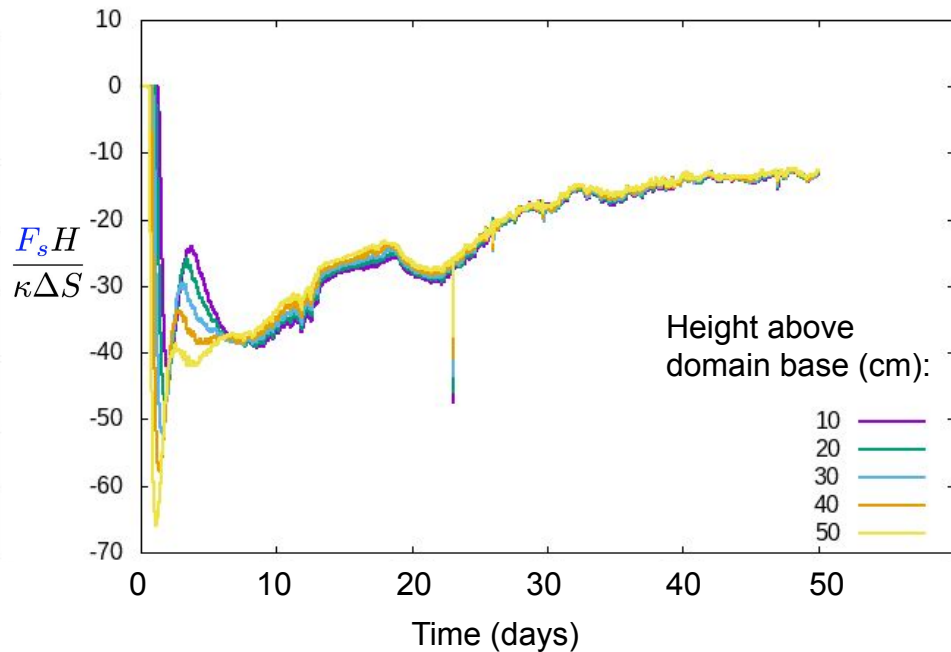
Salt flux weakens in smaller channels as mega-channel develops

Liquid-region salinity and salt flux

Salinity profile evolution, 4x2zbaseline 256³ single-level run



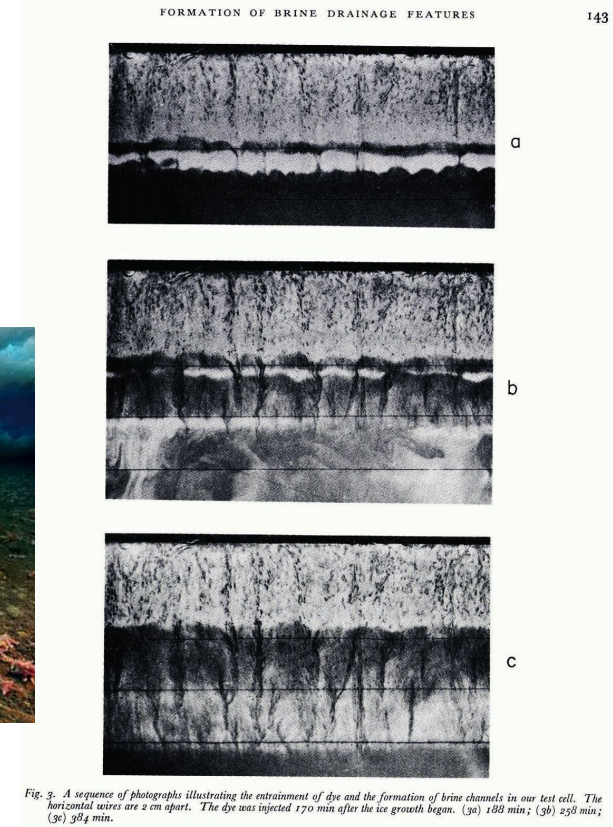
Vertical salinity Flux evolution, 4x4x1 baseline 256x256x64 3-level run



After strong initial desalination pulse, salt flux weakens over time

Discussion

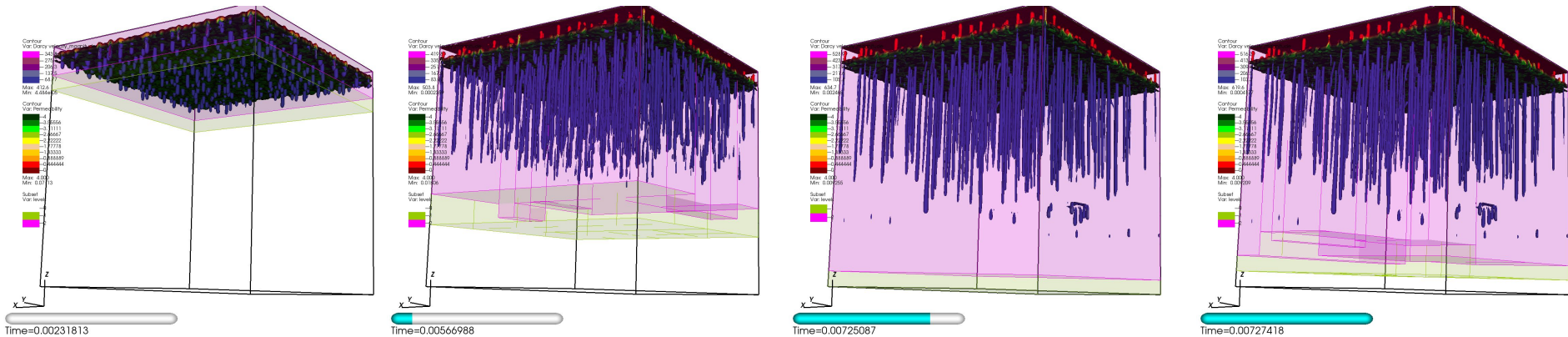
- Initially many small brine channels form, then are consolidated into a single “mega-channel”
 - Single channel is robust over a range of domain sizes
- Shallow region near ice-ocean interface is desalinated by an array of many small brine channels
- Full-depth “mega-channels” allow drainage from saline layer near top of ice
- Comparison with observations:
 - Laboratory: Eide and Martin (1975),
 - “Icy fingers of death” -- BBC Earth



Adaptive mesh refinement

Adaptive mesh refinement focuses computational effort where needed to resolve the problem while using lower resolution in less-dynamic regions.

Initial results are promising but more work needed to fine-tune mesh-refinement criteria



Shaded regions show refinement. Clear is base resolution, green is 2x finer, and purple is 2x even finer (4x base resolution). Over-aggressive refinement in early phases leads to refining the entire domain and slows computation.

Conclusions

- We have simulated brine drainage via 3-dimensional convection in porous mushy sea ice
- Shallow region near ice-ocean interface is desalinated by an array of many small brine channels
- Full-depth “mega-channels” allow drainage from saline layer near top of ice
- Adaptive mesh refinement capability is implemented and is being fine-tuned.

Reference: J.R.G. Parkinson, D.F. Martin, A.J. Wells, R.F. Katz, “Modelling binary alloy solidification with adaptive mesh refinement”, *Journal of Computational Physics: X*, Volume 5, 2020, <https://www.sciencedirect.com/science/article/pii/S2590055219300599>

Appendix: Governing Equations

Continuous equations for conservation of momentum (1), mass (2), salt (3) and energy (4) are found by averaging over lengths greater than the pore scale of sea ice [4, 5].

$$1. \quad \vec{U} = -\frac{k_w(\chi)}{\eta}(\nabla p - \rho_l \vec{g})$$

$$2. \quad \nabla \cdot \vec{U} = 0$$

$$3. \quad \frac{\partial S}{\partial t} = \vec{U} \cdot \nabla S_l = \nabla \cdot \chi D_l \nabla S_l$$

$$4. \quad \frac{\partial H}{\partial t} + \rho_0 c_{p,l} \vec{U} \cdot \nabla T = \nabla \cdot [k_l \chi + (1 - \chi)k_s] \nabla T$$

\vec{U} (Darcy Velocity), χ (porosity), p (pressure), T (temperature),
 S_l (liquid salinity), $S = \chi S_l$ (bulk salinity), η (viscosity), D_l (salt diffusivity),
 α, β (thermal, haline expansion), $c_{p,l}, c_{p,s}$ (liquid/solid specific heat),
 k_l, k_s (liquid, solid heat conductivity), K_0 (reference permeability)

$$H = \rho_0 \{L\chi + [\chi c_{p,l} + (1 - \chi)c_{p,s}]T\} \text{ (enthalpy)}$$

$$\rho_l = \rho_0 [1 - \alpha T + \beta S_l] \text{ (liquid density)}$$

$$K(\chi)^{-1} = (d^2/12)^{-1} + [K_0 \chi^3 / (1 - \chi)^2]^{-1} \text{ (permeability)}$$

Appendix: Computational Approach

Solve (1)-(4) using Chombo finite volume toolkit:

- Momentum and mass: projection method [3].
- Energy and solute:
 - Advective terms: explicit, 2nd order unsplit Godunov method.
 - Nonlinear diffusive terms: semi implicit, geometric multigrid.
 - Timestepping: 2nd order Runge-Kutta method. *Twizell, Gumel, and Arigu (1996)*.

Reference:

James R.G. Parkinson, Daniel F. Martin, Andrew J. Wells, Richard F. Katz, “Modelling binary alloy solidification with adaptive mesh refinement”, *Journal of Computational Physics: X*, Volume 5, 2020, <https://www.sciencedirect.com/science/article/pii/S2590055219300599>